# UPPER ORDOVICIAN TRILOBITES FROM THE ZAP VALLEY, SOUTH-EAST TURKEY 

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#### Abstract

In the extreme south-east of Turkey the Şort Tepe Formation rests disconformably on Arenig strata of the highest Seydişehir Formation. Trilobites are described from the stratotype of the Şort Tepe Formation and a supplementary section, both on the north-east side of the Zap Valley, 40 km south-southwest of Hakkâri. The twenty taxa described include three new species: Sinocybele? fluminis, Calymenesun longinasuta, and Paraplillipsinella pilula. The assemblages, of early Ashgill, probably Pusgillian or possibly Cautleyan, age, consist mainly of genera widespread in Europe, Scandinavia, and Asia, but Sinocybele?, Calymenesun, and Paraphillipsinella have not previously been recorded outside China.


THE south-eastern corner of Anatolia, known in antiquity as the Hakkâri, is a mountainous, traditionally isolated region, bounded to the east by Iran and to the south by Iraq (Text-fig. 1). It is traversed by the River Zap (known sometimes as the Great Zab) which rises near the Iranian border, flows south-west and then south in the vicinity of the towns of Hakkâri and Çukurca, and crosses into Iraq where it forms a tributary of the River Tigris.

Between Hakkâri and Çukurca the Zap cuts a deep valley to expose two inliers of Cambrian and Ordovician sediments, mostly clastics, that form part of the Arab Platform. Until recently the only available account of the Lower Palaeozoic strata was that of Altinli (1963, pp. 60-61), who divided the pre-Devonian rocks into two parts, of Cambrian(?) and Silurian(?) age; the latter, over 1000 m thick and termed Giri Formation, were said to contain Cruziana, possibly of Ordovician age. The latter record led Dean (1980, p. 7) to suggest that, by comparison with the Taurus Mountains, the term Seydişehir Formation should be used in place of Giri Formation. Following initial reconnaissance work, the region was re-mapped by the Turkish Petroleum Corporation (TPAO), whose maps formed the basis of a reassessment of the Cambrian and Ordovician stratigraphy by Dean et al. (1981). The Lower Palaeozoic rocks were noted briefly by Janvier et al. (1984, p. 148, fig. 1), whose account included TPAO's map showing the two inliers to represent the cores of east-west folds, the larger, northerly one termed the Zap anticline, and the other the Çukurca anticline. In Dean et al. (1981) it was demonstrated that shales and sandstones of Upper Cambrian and lower Ordovician age represented the Seydişehir Formation, described from the western Taurus Mountains but widespread in the eastern Taurus, south-eastern Turkey, and neighbouring parts of Iraq. Disconformably overlying strata, mainly shale, mudstone, and quartzite with very minor limestone, were named the Şort Tepe Formation and shown to be of Ashgill age.

## STRATIGRAPHY AND FOSSIL LOCALITIES

## a. Section at Şort Tepe.

The type section of the Şort Tepe Formation is located at the eponymous hill, 7.5 km north-west of Çukurca, where the disconformable junction with the underlying Seydişehir Formation is exposed high on the north-east side of the Zap Valley (text-fig. 1). The line of contact is sharp and planar, with no irregularity or conglomerate at the base of bed a (text-fig. 2), 30 cm of grey oolitic limestone in which no macrofossils were found. Bed b, grey shale 1 m thick, proved sparsely fossiliferous, yielding only a few small brachiopods (Aegiromena? sp.) and fragments of diplograptid graptolites, but no trilobites. No macrofossils of any kind were found in the 1.5 m of bioturbated,

text-fig. 1. Left: sketch map of south-eastern Turkey, showing principal place-namcs. Right: geological map (after TPAO) of the Zap Valley between Hakkâri and Çukurca. $1=$ Zabuk and Sadan formations, undifferentiated (Cambrian); $2=$ Koruk, Seydişehir, and Şort Tepe formations, undifferentiated (Cambrian, Ordovician); $3=$ Upper Palaeozoic rocks (undifferentiated); $4=$ Mesozoic and Tertiary rocks (undifferentiated); Thr $=$ thrust; Flt $=$ fault.
grey siltstone that make up bed c. The most varied assemblages at this section came from bed d, grey shale 5 m thick in which trilobites were found at two levels (locs. Z.33-3, Z.33-4), though much less abundantly than at Şort Dere.

The 2 m of grey-green siltstone of bed e mark a transition from the shale of bed d to a succession of resistant quartzites, mostly thickly bedded but some finely laminated, that are grey-green when fresh but weather to form a distinctive, whitish feature in the hillside. In the higher part of the measured section a small fault, with downthrow to the north-east, cuts the quartzites, whose outcrop continues north-westwards to the adjacent road, where a succession 25 m thick was seen. Silurian rocks are unknown from the area and Devonian strata overlie both Seydişehir and Şort Tepe formations with low angular unconformity in the vicinity of Köprülü, north-west of Şort Tepe (text-fig. 1; see also Janvier et al., 1984, pp. 149-151).
Fannal lists: Z.33-1. Aegiromena? sp., fragments of diplograptid graptolites; Z.33-3. Lonchodomas sp., Dindymene? sp., Prionocheilus cf. obtusus; Z.33-4. Dindymene? sp., Calymenesun altinasuta, Birmanites latus.

## b. Section at Şort Dere

About 1.5 km south-east of Şort Tepe the small valley of Şort Dere intersects the east bank of the River Zap. Approximately 200 m north-east of the intersection, the base of the Şort Tepe Formation rests disconformably on silty shale and quartzite of the Seydişehir Formation, 40 m of which were seen between this point and the Zap. The succession (text-fig. 2), which differs in detail from that at Şort Tepe, is more accessible and better exposed but much less complete and so was used to supplement the stratotype. Again, the interformational boundary is planar, but in this case the basal unit, bed $\mathrm{a}^{\prime}$, is a high energy grainstone, 75 cm thick, ferruginous and oolitic, containing reworked quartzitic fragments derived from the Seydişehir Formation (Monod in Dean et al. 1981, p. 277). No macrofossils were found in the two succeeding units, beds $b^{\prime}$ and $c^{\prime}$, comprising,
text-fig. 2. Measured sections in the Şort Tepe Formation showing stratigraphic positions of fossil localities.

respectively, $1 \cdot 1 \mathrm{~m}$ of grey shale and 40 cm of sandy limestone, but bed $\mathrm{d}^{\prime}$ yielded numerous specimens at two levels, Z. 34 and Z.36. The rocks are mainly brown-weathering, green-grey shale and silty mudstone, often exhibiting bioturbation; 8 m were seen and the highest beds are faulted against Seydişehir Formation, though the contact is masked by rock falls. Most of the fossils were found at loc. Z.34, the great majority of them trilobites, with a few poorly preserved brachiopods and sporadic machaeridian plates and echinoderm debris. All but one (Pl. 62, fig. 8) of the trilobites were disarticulated, and most came from a layer 5 cm thick, though Z. 34 is taken here to include also material from the overlying 40 cm of sediment. The assemblage proved to be diverse, and the asaphinid B. latus, the largest form present, was easily the most abundant.
Fautal lists: Z.34. Ampyx? sp., Lonchodomas sp., Raphiophorns? sp., Hibbertia sp., Sinocybele? fluminis, Ovalocephalus tetrasulcatus, Duftomia sp., C. longinasuta, P. cf. obtusus, Paraphillipsinella pilula, B. latus, Harpidella sp., Phorocephala sp., Amphitryon? sp., Lichas aff. laciniatus, Dicranopeltis sp., Diacanthaspis sp., Miraspis sp.; Z.36. Lonchodomas sp., Dindymene? sp., Paraphillipsinella pilula, Amphitryon? sp., Stenopareia sp.

## AGE AND RELATIONSHIPS OF THE TRILOBITES

During recent years it has become increasingly apparent that many Ordovician trilobite genera once thought to have a limited vertical distribution within the system have, in fact, very long ranges, and that their lateral distribution may be restricted by changes in facies. These comments apply to at least eight trilobites of the Şort Tepe Formation, assessment of whose age depends heavily on only a few genera and species. For present purposes the trilobites are divided into three groups: $a$, genera and species restricted to the Ashgill Series, though sometimes widely distributed;
$b$, genera with a longer stratigraphic range but previously reported only from China; and $c$, genera with both long stratigraphic range and wide geographic distribution.
Group a. Duftonia was described first from the Pusgillian Stage, lowest Ashgill Series, of northern England but has been recorded from the higher Ashgill (pre-Hirnantian) in Wales and Bohemia (Králův Dvůr Formation), though not from the Caradoc or the Silurian.

Birmanites as now interpreted (see Zhou et al. 1984, p. 17 for synonymy) has a long vertical range, from Tremadoc to Ashgill. B. latus is known only from the Ashgill of Västergötland, Sweden, where it occurs in the Red Tretaspis Mudstones, strata equated by Kielan (1960, p. 78) with the combined zones of Eodindymene pulchra and of Staurocephalus clavifrons in the lower and middle portions of the Polish Ashgill. In Scania, southern Sweden, a zone of Opsimasaphus (now B.) latus and Dicellograptus complanatus was employed by Glimberg (1961, p. 83). Jaanusson (1963, pp. 163, 164) equated the O. latus Zone with the E.pulchra Zone, and the D. complanatus Zone with the lower half of his Jerrestad Stage (subsequently termed Jerrestadian, Jaanusson 1982, p. 8), underlain by the Pleurograptus linearis Zone.

Ovalocephalus tetrasulcatus was described (as Hambatocnemis) from the S. clavifrons Zone of the Ashgill in Poland (Kielan 1960, p. 141) and has not been reported elsewhere in Europe. The record from the Şort Tepe Formation increases considerably the geographic distribution of the species, but the genus already had an extended history in China, where it occurs from the Arenig to the Ashgill (Lu and Zhou 1979).

Although the Turkish specimens of Lichas are specifically undeterminable, species assigned to the genus by Tripp (1958, p. 575) occur only in the Ashgill and the lower and middle Silurian. Dicranopeltis is not recorded below the Ashgill, in which it is poorly represented, and also includes several species from the Middle and Upper Silurian (Tripp 1958, p. 575).
Group b. Three genera are of particular interest as they are unrecorded from Europe but are well represented in China where, however, their vertical range extends far below the Ashgill. The type species of Calymenesun came from the Shihtzupu Formation (Llandeilo Series) of Guizhou Province, but the genus is recorded also from low in the Ashgill (Zhou et al. 1984, p. 29).

Paraphillipsinella (see review in Zhou and Dean 1986, p. 767) was founded on material from the Caradoc of Sichuan Province but occurs also in the lower Ashgill of the Yangtze region.

The generic position of trilobites here termed Sinocybele? is uncertain but closest comparison is with Chinese species, all of which have three pairs of pygidial pleurae in contrast to four pairs in European and Scandinavian species of the possibly related Atractopyge.

Whether Amphitryon? should also be included in group b is debatable. The genus is recorded from the higher Caradoc and the Ashgill Series in Europe, but species in which the preglabellar field of the cranidium is triangular in plan have been described only from China, where the character occurs much earlier, in the Llanvirn, and may merit generic recognition.
Group c. Remaining genera in the Şort Tepe Formation contribute no precise evidence of age. Of the Raphiophoridae, Ampyx and Lonchodomas extend from Arenig to Ashgill. The range of Raphiophorus is uncertain; the type species came from the Black Tretaspis Shale, approximately late Caradoc, of Sweden but the genus is well represented in the Ashgill and Silurian (species listed by Thomas 1978, p. 53). A supposedly Arenig species was excluded by Thomas, and Raphiophorus sp. from the Meadowtown Beds (upper Llandeilo or lowest Caradoc) of Shropshire (Whittard 1955, p. 23, pl. 2, figs. 13-16) comprises poorly preserved meraspids of uncertain position. Hibbertia occurs in both Caradoc and Ashgill, and dindymeninids have a long range within the Ordovician. Harpidella ranges from Ashgill to lower Devonian (Thomas and Owens 1978, p. 72), and species of Phorocephala that lack a preglabellar field are reported from both Caradoc and Ashgill (Zhou and Dean 1986, p. 751). The single cranidium of Stenopareia sp . is of little significance and Prionocheilus cf. obtusus, though closely resembling an Ashgill species from Britain and Sweden, represents a genus that changed relatively little during the Ordovician and whose oldest representatives occur in the lower Arenig of southern France (Dean 1966, p. 300). Odontopleurids constitute
only a minor element in the Şort Tepe assemblages and Diacanthaspis is known from both Caradoc and Ashgill strata; Miraspis has a long range, from lower Ordovician to upper Silurian in Europe (Bruton 1968, p. 42).

To summarize, evidence given in group a favours a lower Ashgill age, corresponding to the lower half of the Jerrestadian Stage in terms of the Swedish succession and the Dicellograptus complauatus Zone in terms of the standard British graptolite zones. Correlation of the Ashgill stages with corresponding graptolite zones is imprecise, but according to Williams et al. (1972) the D. complanatus Zone falls within approximately the upper half of the Pusgillian, though no distinctively Pusgillian trilobites were found in the Şort Tepe assemblages. Although B. latus and O. tetrasulcatus occur together in the Şort Tepe Formation, the latter species is found slightly higher (Stauroceplualus clavifrous Zone) in Poland, so it is possible the Turkish strata may extend above the D. complanatus Zone, into the Cautleyan Stage.

The relationship of the Şort Tepe Formation to successions elsewhere in south-eastern Turkey is not yet established, but the rocks may represent a continuation of the transgressive sequence, represented by the Bedinan Formation, that began in the middle Caradoc and persisted probably into the Ashgill in the Derik-Mardin region, 320 km west of Çukurca (Dean et al. 1981, p. 278).

## SYSTEMATIC DESCRIPTIONS

Terminology is cssentially that used in the Treatise on Invertebrate Paleontology (Harrington et al. in Moore 1959, pp. O117-O126), with the addition of eye socle (Shaw and Ormiston 1964) and baccula (Öpik 1967). Stratigraphic position of fossil localitics at Şort Tepe and Şort Dere is shown in text-fig. 2. Figured and cited specimens are deposited in the Department of Palaeontology, British Museum (Natural History), London, and their numbers carry the prefix It.

Family raphiophoridae Angelin 1854
Genus ampyx Dalman 1827
Type species. Anupyx hasutus Dalman 1827, Asaplus Limestone (uppcr Arenig) of Västanå, Östergötland, Sweden.

## Aupyx sp.

Plate 58 , figs. 2?, 8?, 9
Figured specimens. It. 19494 (Pl. 58, fig. 2), It. 19497 (Pl. 58, fig. 8), It. 19499 (Pl. 58, fig. 9).
Locality. Şort Dere, Z.34.
Description and discussion. The pygidium resembles that of $A$. nasutus, refigured by Whittington (1950, pl. 74, figs. 3-9), but is slightly shorter and the anterior pleural furrows are almost straight. The cranidium is slightly crushed but the shape of the glabella is generally similar to that of $A$. nasutus in having three pairs of depressed muscle scars on the glabellar flanks, and in the form of the anterior border, defined by a shallow border furrow that is continuous with the axial furrows. It differs from $A$. nasutus in having the anterior branches of the facial suture less convergent forwards. In all these characters the Turkish species is comparable with A. abnormalis Yi (1957, p. 557, pl. 5, fig. 3a-e; Lu 1975, p. 414, pl. 39, figs. 5-11; pl. 40, figs. 1-7), from the upper Arenig to Llandeilo of western Hubei, China, but the pygidium of the latter has a narrower, more gently tapered axis and the pleural regions have four pairs of pleural furrows. Rccords of Anpyx from Ashgill strata are rare and the Turkish matcrial may represent a new spccies, but is too poorly preserved for formal description.

Genus lonchodomas Angelin 1854
Type species. Ampyx rostratus Sars 1835, Ampyx Limestone, $4 \mathrm{a} \beta$ (Llandeilo or lowest Caradoc) of Bygdoy, Oslo, Norway.

## Lonchodomas sp.

Plate 58, figs. 5, 6, 11
Figured specimens. It. 19495 (Pl. 58, fig. 5), It. 19496 (Pl. 58, fig. 6), It. 19498 (Pl. 58, fig. 11).
Localities. Şort Tepc, Z.33-3; Şort Dcre, Z. 34 and Z.36.
Description and discussion. The incomplete, slightly compressed cranidium is characterized by the wide (tr.) triangular outline, the short (sag.) anterior projection of the glabella, and the slight curvature, abaxially concave, of the axial furrows. These features suggest comparison with L. tecturmasi (Weber 1932, p. 6, pl. 4, fig. 43; 1948, p. 18, pl. 2, figs. 20-22, 26; Chugaeva 1958, p. 32, pl. 2, figs. 3-5) from the Anderken Horizon (Caradoc) of the Chu-Ili Mountains, Kazakhstan and L. jiantsaokouensis Lu (1975), p. 421, pl. 41, figs. 11 and 12) from the Jiantsaokou Formation (low Ashgill) of northern Guizhou, China. But the Turkish specimen is inadequate for refercnce to either of these species.
L. portlocki (Barrande 1846, p. 9; 1852, p. 636, pl. 30, figs. 24-28; Olin 1906, p. 69, pl. 4, figs. 5-8; Kielan 1960 , p. 169 , pl. 33, fig. 8; pl. 35, fig. 4), from the Ashgill of Bohemia, Sweden, and Poland, also bears some resemblance to the Turkish form, but is distinguished by its narrower (tr.) cranidium and what appears to be a slightly depressed preoccipital lobe.

## Genus Raphiophorus Angelin 1854

Type species. Raplioplorus setirostris Angelin 1854, Lower Tretaspis Shale (Ashgill) of Draggå bro, Dalarne, Sweden.

## Raphiophorus? sp.

## Plate 58, figs. 1, 3, 4, 7

Figured specimens. It. 19490 (Pl. 58, fig. 1), It. 19491 (Pl. 58, fig. 3), It. 19492 (Pl. 58, fig. 4), It. 19493 (Pl. 58, fig. 7)

Locality. Şort Dere, Z.34.
Description and discussion. The glabella is similar to that of R. setirostris Angelin (see Whittington 1950, p. 553, pl. 74, figs. 1 and 2), a species in which, according to Whittington, triangular bacculae are probably present, as they are in the Turkish material. However, the wide (tr.), long (exsag.) fixigenae, the forward curvature of the distal portions of posterior border and furrow, and the short anterior projection of the glabella in front of the fixigenae distinguish the Turkish specimens from R. setirostris. These features suggest, rather, a comparison with the type species of Taklamakania, T. tarimensis W. Zhang (1979, p. 1003, pl. 1, fig. 9), from the Engou Formation (Caradoc) of Keping, Xinjiang, China. A specifically identical specimen from the same locality and horizon was later described as a new genus and species Xinjiangia yinganensis T. Zhang (1981, p. 199, pl. 74, fig. 11). However, Taklamakania has a larger, longer pygidium than Rapliophorus and its thorax comprises only three segments. The pygidium of the Turkish species generally resembles that

## EXPLANATION OF PLATE 58

Figs. 1, 3, 4, 7. Raphiophorus? sp. Loc. Z.34. 1, cranidium, It. 19490, $\times 5 \cdot 5$. 3, cranidium, It. 19491, $\times 6$. 4, cranidium, It. 19492, $\times 5 \cdot 5$. 7, pygidium, It. 19493, $\times 8$.
Figs. 2 and 8. Ampyx? sp. 2, cranidium, It. 19494, $\times 5$, loc. Z.36. 8, pygidium, It. 19497, $\times 5$, loc. Z. 34 .
Figs. 5, 6, 11. Lohchodomas sp. 5, cranidium, It. 19495, $\times 6$, loc. Z.34. 6, pygidium, It. 19496, $\times 4$, loc. Z.33F.3. 11, hypostoma, It. 19498, $\times 6$, loc. Z.36.

Fig. 9. Ampyx sp. Loc. Z.34. Cranidium, It 19499, $\times 5$.
Figs. 10, 12, 13, 17. Hibbertia sp. Loc. Z.34. 10, dorsal surface of left genal prolongation, It. 19500, $\times 3 \cdot 5$. 12, ventral surface of left genal prolongation, It. 19501, $\times 4$. 13, ventral surface of left genal prolongation, It. 19502, $\times 4.17$, part of right genal region of cranidium, It. 19503, $\times 3 \cdot 5$.
Figs. 14-16, 18. Dindymene? sp. 14, cranidium, It. 19504, $\times 5 \cdot 5$, loc. Z.33-4. 15, cranidium, It. 19505, $\times 6$, loc. Z.36. 16, cranidium, It. 19506, $\times 5$, loc. Z.33-3. 18 , cranidium, It. 19507, $\times 6$, loc. Z.33-3.

of Raphiophorus but has three distinct axial rings, traces of a fourth ring, and three pairs of wide (exsag.) pleural furrows. In $R$. setirostris there are two axial rings and one well-defined pair of pleural furrows.

Family harpetidae Hawle and Corda 1847
Genus hibbertia Jones and Woodward 1898
Type species. Harpes flanaganni Portlock 1843, Bardahessiagh Beds (early Caradoe) of Pomeroy, County Tyrone, Northern lreland.

Hibbertia sp.
Plate 58, figs. 10, 12, 13, 17; Plate 59, fig. I
Figured specimens. It. 19500 (Pl. 58, fig. 10), It. 19501 (Pl. 58, fig. 12), It. 19502 (Pl. 58, fig. 13), It. 19503 (Pl. 58, fig. 17), It. 19508 (Pl. 59, fig. 1).

Locality. Şort Dere, Z. 34.
Description and discussion. The speeimens exhibit a uniformly very wide brim with large genal prolongations that narrow gently posteriorly, features suggestive of Hibbertia. The cheek roll is narrow, widens slightly medially, and is well defined by the distinct girder, which dies out before reaching the posterior margin. The remains of the glabella suggest that it was slightly pointed frontally. The eye ridge is narrow and the ala depressed, defined by a distinct alar furrow as in H. sanctacrucensis Kielan (I960, p. 157, pl. 34, figs. 4 and 6; pl. 35, fig. 8), from the Ashgill Series, S. clavifrons Zone, of Brzezinski, Poland. The fringe is finely and densely pitted with small pits of almost uniform size, and the cheek is eovered with radiating, anastomosing ridges.

## Family encrinuridae Angelin 1854 <br> Subfamily cybelinae Holliday 1942 <br> Genus sinocybele Sheng 1974

Type species. Sinocybele baoslanensis Sheng 1974, Lower Pupiao Formation (Llandeilo to Caradoc Series), south of Shihtien, western Yunnan, China.

Sinocybyle? fluminis sp. nov.
Plate 59, figs. 2-6, 7?, 8, 9
Diagnosis. Sinocybele? species with four pairs of large tubercles sub-equispaced along median area of glabella. Palpebral lobes sited opposite posterior half of 2 p glabellar lobes and about midway between glabella and lateral border furrow. Posterior branches of facial suture are very slightly curved, convex forwards, and meet the margin just in front of genal angles. Strongly developed eye ridges end opposite 3p furrows. Pygidium with three pairs pleurae that curve strongly backwards

## EXPLANATION OF PLATE 59

Fig. 1. Hibbertia sp. Loc. Z.34. Ventral surface of cephalic fringe, It. 19508, $\times 3$.
Figs. 2-6, 7?, 8, 9. Sinocybele? fluminis sp. nov. Loc. Z.34. 2, cranidium, It. 19509, $\times 5$. 3, pygidium, It. $19510, \times 5.4$, pygidium, It. $19511, \times 6.5$, pygidium, It. $19512, \times 5.6$, pygidium, It. $19513, \times 5.7$, hypostoma referred questionably to speeies, It. 19514, $\times 6.8$, left librigena with assoeiated pygidium of Paraphillipsinella, 1t. 19515, $\times 5.9$, cranidium, It. 19516, $\times 5.5$ is holotype; remainder (excluding fig. 7) are paratypes.
Figs. 10, 12-16. Ovalocephalus tetrasulcatus (Kielan I960). Loc. Z.34. 10, cranidium, 1t. 19477, $\times 5$. 12, pygidium, It. $19517, \times 5$. 13, cranidium, It. $19478, \times 6.14$, cranidium, 1t. 19518, $\times 6$. 15, hypostoma, It. $19519, \times 5.16$, cranidium, It. $19520, \times 5$.
Figs. 11, 17, 18. Duftonia sp. Loe. Z.34. 11 and 18, internal mould and latex cast of cranidium, It. 19521, $\times 3 \cdot 5.17$, cranidium, It. 19522, $\times 3$.


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to end in long free points. Axis subtriangular with three complete and about nine partially defined axial rings.

Holotype. It. 19512 (pl. 59, fig. 5).
Paratypes. It. 19509 (Pl. 59, fig. 2), It. 19510 (Pl. 59, fig. 3), It. 19511 (Pl. 59, fig. 4), It. 19513 (Pl. 59, fig. 6), It. 19515 (Pl. 59, fig. 8), It. 19516 (Pl. 59, fig. 9).

Locality. Şort Dere, Z. 34 .
Description. The more complete of the two paratype eranidia has an estimated overall breadth (excluding fixigenal spines) of 11.6 mm and an estimated length of 3.7 mm . Even allowing that the specimen is dorsally eompresscd, the long (tr.), acutely triangular outline of the posterior areas of the fixigenae is noteworthy. The combined glabella and preglabellar field is about as broad as long, its outline broadening slightly so that the basal breadth is 0.8 of the frontal. The boundary between glabella and preglabellar field is indieated by a pair of short (tr.), shallow grooves (PI. 59, fig. 2) that run adaxially forwards from the axial furrows. Three inequisized pairs of lateral glabcllar lobes, 1 p the smallest, are separated by short (tr.), deep lateral furrows, the 1 p pair of which turn backwards adaxially so that the 1 p lobes are of 'cat's ear' outline, linked to the median area by narrow necks. Most of the glabella, if not all, is covered with fine granules. In addition there are four pairs of large tubcrcles, subequally spaced longitudinally and sited opposite the rear half of the $2 p$ lobes, the centre of the 3 p lobes, and then at almost equal intervals between the 3 p lobes and the anterior border. The last named is low, well defined by a broad (sag.) anterior border furrow, and carries about six (estimated) large tubercles dorsally; it is incompletely preserved medially and the presence of a median projection has not yet been demonstrated.

The more complete cranidium shows the anterior branehes of the facial suture to be straight, converging forwards slightly from eyes that are sited opposite the lp glabellar furrows and approximately midway between the axial furrows and the lateral border furrow. The cye lobes are joined to the axial furrows opposite the 3p glabellar furrows by strongly dcveloped, smooth eye ridges. One of the latter is better seen in the seeond paratype cranidium (Pl. 59, fig. 9), in which, owing to crushing, it appears to be directed more strongly backwards.

A single left librigena (Pl. 59, fig. 8) shows the eye to be small, possibly pedunculate, and the lateral border well defined; the pitted surface carries a few large tubereles like those on the lateral border.

The pygidium bears a marked general resemblance to that of Atractopyge Hawle and Corda 1847 (for examples, see Dean 1971 and Ingham 1974), the most obvious difference being that in S.? fluminis the pleural regions are composed of three pairs of plcurac instead of four. The outline of the axis is subtriangular, slightly constricted at the third ring furrow, and apparently ends in a sharp point that represents, in fact, the post-axial ridge, set slightly below the parabolie, diminutive, true terminal piece (Pl. 59, figs. 5 and 6 ). The first three ring furrows are complete but subsequent furrows (nine, possibly ten are visible) are incomplete both immediately adjacent to the axial furrows and medially, where a smooth median band oceupies the axial third of the axis. Three pairs of pleurae that end in long free points curve strongly backwards and inwards so that the third pair almost meet behind the axis, where they are separated by the small post-axial ridge. Each pleura is divided by a distinet pleural furrow into unequal anterior and posterior bands, the latter about twiee as wide (tr.) as the former. Anterior bands are clearly visible on the first two pairs of pleurae but searcely or not at all on the third pair, so that the second and third pairs of posterior bands appear to be separated by a single furrow, in which a row of granules may be visible. Surface of the pygidium is granulose.

The hypostoma of $S$. baoshanensis remains undescribed and the present small specimen (Pl. 59, fig. 7) is therefore assigned only questionably to the new species. It is of cncrinurid type and its outline (excluding anterior wings) is suboval, with maximum breadth 2.0 mm . and median length 2.5 mm . The anterior twothirds of the middle body are divided into three longitudinal lobes by a pair of straight, parallel furrows. The central lobe so formed occupies just over half the breadth of the middle body and projects forwards of it; in life position it would have underlain the centre of the anterior border. Temple (1954, p. 318) suggested that in Encrinurus the generally comparable central lobe may have aecommodated the ventral surface of the pygidium during enrolment and the present specimen may have functioned similarly. The middle body is cireumseribed by a narrow, rim-like border that broadens to form the incompletely preserved anterior wings.

Discussion. S. baoshanensis Sheng (1974, p. 110, pl. 7, fig. $6 a, b$ ) was founded on a single cranidium, illustrated as both internal and external moulds, of cybelinid type in which the anterior border
was produced forwards to form a prominent, flat-topped, horn-like protuberance whose length, though incomplete, was at least two-thirds that of the glabella. Three, possibly four, pairs of large tubercles are visible on the central area of the glabella and the remainder of the surface is covered with small tubercles.

Zhou et al. (1984, p. 27), in discussing the relationships of Cybelurus? from the Shihtzupu Formation in Guizhou Province, noted that Sinocybele has branching 3p glabellar furrows like Cybelurus and that if the pygidium of Sinocybele proves to have three pairs of pleurae, their Cybelurus? sp. would be assignable to Sinocybele even though the anterior cranidial projection is only small. A. sinensis Lu (1975, pp. 233, 444, pl. 45, figs. 15 and 16), from the Shihtzupu Formation of Zunyi and based on a single cranidium, was assigned to Cybelurus? by Zhou et al. (1984, p. 27). The holotype has long (tr.), acutely triangular, postocular fixigenae generally similar to those of S. fluminis, but the eyes of sinensis (here also placed questionably in Sinocybele) are set very far back, opposite the 1 p glabellar lobes; both species have the fixigenal spines directed only slightly backwards. $S$.? sinensis has five pairs of large tubercles on the median area of the glabella, and there are additional, paired tubercles on the abaxial parts of the combined frontal glabellar lobe and preglabellar field.

Subfamily dindymeninae Henningsmoen in Moore 1959
Genus dindymene Hawle and Corda 1847
Type species. Dindymene fridericiaugusti Hawle and Corda 1847, from the Králův Dvůr Formation (Ashgill Series), Králův Dvůr, Czechoslovakia.

## Dindymene? sp.

Plate 58, figs. 14-16, 18
Figured specimens. It. 19504 (Pl. 58, fig. 14), It. 19505 (Pl. 58, fig. 15), It. 19506 (Pl. 58, fig. 16), It. 19507 (Pl. 58, fig. 18).

Localities. Şort Tepe, Z.33-3 and Z.33-4; Şort Dere, Z.36.
Description and discussion. The Turkish specimens, inadequate for specific determination, comprise mostly fragments of cranidia, though one piece of shale (1t. 16063, not illustrated) shows vestiges of a partly disarticulated thorax and pygidium of dindymeninid type, the pleural tips extended to form long, slim spines. The posterior border and border furrow are narrow, transversely straight, and the genal angles are produced to form fixigenal spines, seen in It. 19507. The same specimen shows a broken spine, c. 2 mm long, apparently extending from the glabella, and It. 19505 (Pl. 58, fig. 15) may retain the spine base. An analogous though shorter spine is seen on some cranidia of D. lughesiae Reynolds 1894 (Ingham 1974, pl. 18, figs. 1, 4, 8 -10), from the Ashgill (Rawtheyan) of northern England and of D. cordai Nicholson and Etheridge 1878 (1ngham 1974, pl. 18, fig. 18), from the Rawtheyan of Scotland; a spine base is visible on the holotype of $D$. fridericiaugusti (original of Hawle and Corda 1847, pl. 1, fig. 3; see also Horný and Bastl 1970, pl. 15, fig. 1). The remainder of the glabellar surface is smooth except for a few, widely spaced tubercles, some possibly paired. By contrast the fixigenal surface is corsely pitted, with more numerous large tubercles than on the glabella. Broadly similar ornamentation is seen on cranidia of D. ornata Linnarsson 1869 from Sweden and Poland illustratcd by Kielan (1960, pl. 26, fig. 6; pl. 27, fig. 4), though the glabella is more granulose and has more tubcrcles than the Turkish species.

## Family hammatocnemidae Kielan 1960 <br> Genus ovalocephalus Koroleva 1959

Type species. Ovalocephalus kelleri Koroleva 1959, from the late Caradoc of northern Kazakhstan. The close resemblance of Ovalocephalus and Hammatocnemis was noted by Zhou and Dean (1986, p. 776) and the two genera are considered here as synonyms.

## Ovalocephalus tetrasulcatus (Kielan 1960)

Plate 59, figs. 10, 12-16
Hanmatocnemis tetrasulcatus Kielan 1960, p. 141, pl. 25, fig. 3; pl. 26, figs. 2-4; pl. 27, figs. 6-8; text-figs. 38 and 39.

Figured specimens. It. 19477 (Pl. 59, fig. 10), It. 19478 (Pl. 59, fig. 13), It. 19517 (Pl. 59, fig. 12), It. 19518 (Pl. 59, fig. 14), It. 19519 (Pl. 59, fig. 15), It. 19520 (Pl. 59, fig. 16),

Locality. Şort Dere, Z. 34 .
Description and discussion. All the Turkish specimens are small, but the cranidia match closely those illustrated from Poland. According to Kielan's (1960, p. 142) original account the frontal breadth of the glabella is equal to threc times its breadth in front of the preoccipital segment, but her illustrations of undistorted specimens (Kielan I960, pl. 26, fig. 4; pl. 27, fig. 6; text-fig. 38) show that it is only 2.3 times as broad. In the largest, slightly compressed, Turkish cranidia the corresponding figures are $2 \cdot 2$ and $2 \cdot 4$, and the specimens closely resemble the holotype and one paratype (Kielan 1960, pl. 26, figs. 2 and 4).

The hypostoma of Ovaloceplatus is not well known and that of $O$. tetrasulcatus has not been described, but the present Turkish example (Pl. 59, fig. 15) is attributed to the species on account of its resemblance to the hypostoma of $O$. decorus (Lu in Lu and Chang 1974) figured by Lu and Zhou (1979, pl. 3, fig. 6). The specimen is almost as long ( 2.8 mm ) as broad ( 3.0 mm ), of low convexity, pentagonal in outline with the transverse anterior margin slightly convex. The subparallel lateral margins occupy 0.57 of the overall length, and the straight posterolateral margins converge to meet at an angle of $100^{\circ}$. A low, narrow rim runs around the lateral and posterior margins and widens (sag.) slightly to form a small point at the posterior extremity of the hypostoma.

The only available Turkish pygidium ( Pl .59 , fig. 12) is very small, with an estimated breadth and median length of 4.0 mm and 1.4 mm respectively, and closcly resembles the Polish examplcs. The axis has three distinct axial rings and a fourth is less well defined. In the Polish type material the pleural regions comprise four pairs of pleurae, the first three distinct, and four pairs of free points were said to be present, though these are not seen in all the illustrations. In the Turkish example there are three distinct pleurae plus faint traces of a fourth. First and sccond pleurae are bounded by broad (exsag.), deep, interpleural furrows and end in short free points; third pleurae show no free points and only the adaxial half of the third interpleural furrow is clearly defined. All three pleurae have a node developed immediately outside the axial furrow; a similar structure was described by Kielan (1960, p. 143) and evidently corresponds to nodes on the thorax (see also Lu and Zhou's illustrations I979, pl. 4, figs. 3 and 4 of the thorax of O. decorus (Lu in Lu and Chang 1974)).

Present evidence suggests that $O$. tetrasulcatus has been found as yet only in Poland and south-eastern Turkey. O. tetrasulcatus as recorded by Lu and Zhou (1979, pl. 2, figs. 10 and 11) from the Qilang Formation (Caradoc) of Keping, Xingjiang Province, China, has since been described as O. kanlingensis (T. Zhang 198I, p. 209, pl. 77, figs. 5-7).

## Family dalmanitidae Vogdes 1890 <br> Genus duftonia Dean 1959

Type species. Duftonia lacunosa Dean 1959, Dufton Shales (Ashgill; Pusgillian Stage) of northern England.

## Duftonia sp.

Plate 59 , figs. $11,17,18$
Figured specimens. It. 19521 (Pl. 59, figs. 11 and 18), It. 19522 (Pl. 59, fig. 17).

## Locality. Şort Dcre, Z. 34

Description and discussion. The Turkish cranidia differ from D. lachnosa Dcan (1959, p. 144, pl. 19, figs. 2, $5,6,8$ ) in having: frontal glabellar lobe, though slightly compressed, proportionately longer, greater than half the glabellar length, compared with about half; rear ends of the cye set proportionately further from the axial furrows and opposite the mid-points of the 2 p glabellar lobes, compared with opposite the 1 p glabellar furrows. In the two species both the palpebral lobes and the well-defined, strongly sigmoidal palpebral
furrows extend to the axial furrows, in D. lacurosa at points opposite the $3 p$ glabellar furrows and anterolateral tips of the $3 p$ lobes, and in $D$. sp. well in front of the $3 p$ furrows. The Turkish matcrial shows traces of a very low, thin rim that represents the anterior border, a structure scarcely discernible in $D$. lacunosa.

Dalmanites morrisiana Barrande (1852, p. 559, pl. 27, figs. 8 and 9) was assigned to Duftonia by Dcan (1967a, p. 38) and the original of Barrande's fig. 8 was selected as ncotypc by Marek (in Horny and Bastl 1970, p. 210). Dalmanites morrisiana, from the Králův Dvůr Formation (Ashgill), has slightly smaller, narrower palpebral lobes than Duftonia sp. and these structures extend from just behind the 3p furrows to end at points relatively further from the posterior border furrow than in either $D$. sp. or $D$. lacumosa, and a greater distance from the glabella. Evidence for a median occipital tubcrcle in D. lacrmosa is cquivocal, but one is visible both in D. sp. and in Barrande's illustrations of Dalmanites morrisiana.

Family calymenidae Burmeister 1843
Subfamily reedocalymeninae Hupé 1955
Lu (1975, pp. 445-458) included Reedocalymene, Calymenesum, and Neseuretus (a senior subjective synonym of Symhomalonotus) in the Reedocalymeninae, with the tacit implication that Synhomalonotinac Kobayashi 1960 is a junior synonym. We follow this classification provisionally here as Calymenesun has several characters in common with Neseuretus, and we add Vietnamia Kobayashi 1960 and Neseuretus (Nesenretinus) Dean 1967b. The position of Reedocalymene Kobayashi 1951, with anterior projection of the frontal area still longer than that of Calymenesum, is less clear and the genus is in need of revision.

## Genus calymenesun Kobayashi 1951

Type species. Calymene tingi Sun 1931, Shihtzupu Formation (Llandcilo) of Feilaishi, Zunyi, Guizhou, China.

## Calymenesun longinasuta sp. nov.

Plate 60 , figs. $1-3,5,6,8-10,12,13$
Diagnosis. Calymeuesun species with glabellar outline straight-sided laterally and frontally. Anterior border steeply inclined forwards, well defined by anterior border furrow that is deep abaxially but broad (sag.) and shallow medially. Median third of anterior border of cranidium is produced to form stout spine. Lateral border wide, well defined.

Holotype. It. 19527 (Pl. 60, figs. 8 and 9).
Paratypes. It. 19523 (Pl. 60, figs. 1 and 2), It. 19524 (Pl. 60, fig. 3), It. 19525 (Pl. 60, fig. 5), It. 19526 (Pl. 60. figs. 6, 10, 12, 13).

Locality. Şort Dcre, Z. 34.
Description. The length of the glabella is almost equal to, or slightly less than, its basal breadth; there are three inequisized pairs of lateral lobes, and glabcllar outline tapers evenly to a transversely straight anterior margin. Anterior border is produced in same plane to form a frontal spine at least as long as the preglabellar field. Pedunculate palpebral lobes stand higher than glabella, and are sited opposite the $2 p$ furrows and $3 p$ lobes. Weakly developed eye ridges extend to the axial furrows opposite, or slightly in front of, the $3 p$ furrows. Anterior branches of facial suture are straight and convergent. Axial furrows widen abaxially opposite the 1 p lobes to accommodate a pair of small bacculae. Median occipital tubercle present. Surface, excluding furrows, is mostly granulose but the median lobe of the glabella carries five equispaced pairs of tubercles that become progressively larger from front to rear. Glabella generally resembles that of Neseuretus but the preglabellar field and anterior border are clearly defined, quite apart from the striking development of the anterior spine. The large, paired glabellar tubercles are particularly distinctive and the rearmost pair is visible also on the internal mould.

Paratype left librigena is of typical calymenid form but lateral border is very wide and well defined. The specimen shows the cye surface, though incomplete, to bc short (exsag.), bounded by poorly defined eye socle.

The pygidium is of calymenid type, with seven axial rings and five or six pairs of furrowed ribs. Outline of axis is slightly constricted behind sixth axial ring, and postaxial ridge is apparently parallel-sided and convex as in Neseuretus. Pleural furrows become progressively less well defined from front to rear and do not quite attain the lateral margin. All the ribs are divided by faint interpleural furrows into two unequal bands, the anterior twice as wide (exsag.) as the posterior. About midway between axial furrow and lateral margin is a faint depression that corresponds to what Campbell (1967) termed a cincture, a coaptative structure commonly developed in calymenids.

Discussion. C. tingi (Sun 1931, p. 29, pl. 3, fig. 9a-g only, non 9hl) was redescribed by Zhou et al. (1984, p. 29, fig. $7 a-g, i, j$ ) and differs from the new species in several respects: the glabella widens considerably across the 1 p lobes, and the axial furrows, which contain bacculae, are strongly curved, abaxially concave, in a manner recalling that in Vietnamia Kobayashi (1960, p. 43); the anterior branches of the facial suture are curved; the anterior border is less distinctly defined and forms a process that extends to produce a slim spine as long as the remainder of the cranidium; the eyes are sited opposite the 2 p lobes and furrows; the surface is finely granulose with no tubercles.
C. granulosa Lu (1975, pp. 238, 450, pl. 47, figs. 1-5), from the top of the Linhsiang Formation (lowest Ashgill, Nankinolitlus Zone) at Chikangpo, Ichang district, west Hupei, China, has a proportionately shorter, broader (esp. basally) glabella than the new species; the anterior border and furrow are almost undefined; and the lateral border furrow is absent. In these respects $C$. granulosa is more comparable with C. tingi.

In C. yinganensis Zhang (1981, p. 211 , pl. 78, figs. 3-5), a species previously referred to Neseuretus (Zhang et al. 1982, p. 72, table 10), from the Qilang Formation (Caradoc) of Kanling, Keping, Xinjiang Province, China, the preglabellar field and anterior border, though incomplete, appear less well defined than in the new species; they and the anterior branches of the facial suture are more comparable with those of C. tingi, though the latter is readily recognized by the distinctly large basal breadth of the glabella. Small, sparse tubercles with a suggestion of arrangement in transverse rows ornament the glabella of C. yinganensis, and the pleural regions of the pygidium show well-developed cinctures like those of $C$. tingi, but unlike the new species. C. zhejiangensis Ju in Qiu et al. (1983, p. 250, pl. 87, figs. 11 and 12), from the Huangnekhan Formation (Ashgill, Nankinolithus Zone) of Jiangshan, west Zhejiang Province, China, has distinct bacculae and the uniformly tapered glabellar outline is more like that of C. altinasuta than that of C. tingi, though the latter's less well-defined preglabellar field is more comparable. The pygidium of $C$. zliejiangensis is very different from that of the new species in having a broad (exsag.), deep cincture that divides the pleural regions into small, coarsely ribbed proximal and weakly ribbed distal portions, and coincides with the junction of terminal piece and post-axial ridge.

## EXPLANATION OF PLATE 60

Figs. 1-3, 5, 6, 8 10, 12, 13. Calymenesun Ionginasuta sp. nov. Loc. Z.34. 1 and 2, cranidium, It. 19523, $\times$ 3. 3, pygidium, It. 19524, $\times 4$. 5, left librigena, It. 19525, $\times 3.6$ (internal mould), 10, 12, 13 (latex cast), cranidium, It. 19526, $\times 3.8$ and 9 , cranidium, It. 19527, $\times 4.8$ and 9 , holotype; remainder are paratypes.
Figs. 4 and 24. Plorocephala sp. Loc. Z.34. 4, pygidium, It. 19528, $\times 9$. 24, cranidium, It. 19529, $\times 5$.
Figs. 7, 17-21, 23. Paraphillipsinella pilula sp. nov. 19 from Loc. Z.36; remainder from Loc. Z.34. 7, pygidium, It. 19530, $\times 8$. 17, cranidium, It. 19531, $\times 6$. 18, cranidium, It. 19532, $\times 6.19$, pygidium, It. $19533, \times 8$. 20, cranidium, It. 19534, $\times 6.21$, cranidium, It. 19535, $\times 6.23$, cranidium, It. 19536, $\times 6$. 17 is holotype; remainder are paratypes.
Figs. 11, 14-16. Prionocheilus cf. obtusus (M‘Coy 1846). Loc. Z.34. 11, pygidium, It. 19537, $\times 5$. 14, cranidium, It. 19538, $\times 5.15$, cranidium, It. 19539, $\times 6.16$, pygidium and two, possibly three, attached thoracic segments, It. 19540, $\times 4$.
Fig. 22. Harpidella sp. Loc. Z.34. Cranidium, It. 19541, $\times 6$.


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## Subfamily pharostomatinae Hupé 1953 Genus prionocheilus Rouault 1847

Type species. Prionocheilus verneuili Rouault 1847, from an unnamed formation of Middle Ordovician age at Poligné, Brittany, France.

Prionocheilus cf. obtusus (M‘Coy 1846)

## Plate 60, figs. 11, 14-16

Figured specimens. It. 19537 (Pl. 60, fig. 11), It. 19538 (Pl. 60, fig. 14), It. 19539 (Pl. 60, fig. 15), It. 19540 (Pl. 60, fig. 16).

Locality. Şort Dere, Z. 34.
Description and discussion. M‘Coy’s (1846, p. 54, pl. 4, fig. 6) holotype cranidium from the Chair of Kildare Limestone in eastern Ireland was redescribed by Whittington (1965, p. 55, pl. 16, figs. 1-3, 6) and by Dean (1971, p. 42, pl. 18, figs. $10,12,13$ ), who figured additional topotype material both of the species and of Calymene leptaenarum Törnquist 1884, placed in synonymy with it. The age of the type material is Ashgill, probably Rawtheyan Stage. Both Turkish cranidia are slightly distorted and incomplete but generally resemble the Irish material, and the anterior border and preglabellar field are similar. In both specimens the eye ridges meet the axial furrows just in front of the 2 p furrows as in the holotype of $P$. obtusus, but whether the palpebral lobes are sited opposite the outer ends of the same furrows is less clear. Plate 60 , fig. 14 shows the characteristic widening of the axial furrows opposite the 1 p lobes to accommodate structures that have been termed, variously, Plarostoma-Flecke (Öpik 1937, p. 23) or paraglabellar areas (Harrington et al. in Moore 1959, p. O123).

In Plate 60, fig. 11 the pygidium is slightly shortened by compression but the anterior half of the axis carries four axial rings, followed by traces of a fifth; the remaining pleural region has five well-defined ribs with faint interpleural furrows, and part of a sixth rib in addition to the anterior half rib. Pygidia of $P$. obtusus from Ireland (Dean 1971, pl. 18, figs. 4, 5, 14; pl. 19, figs. 5, 10, 12) and Sweden (Warburg 1925, p. 157, pl. 4, figs. 2-4) have five or six axial rings, the last poorly defined, and five pairs of ribs. The apparently rounded tips of the ribs in the Turkish specimen are due to weathering, and are not an original feature.

The distribution of P. obtusus is not known in detail. Although the type material is probably of Rawtheyan age, and $P$. cf. obtusus from the Rhiwlas Limestone of North Wales is likely to be of similar age, it is clear that broadly comparable forms have an extended stratigraphic range and specimens from the Caradoc of Norway described by Owen and Bruton (1980, p. 32, pl. 9, figs. 10, 11, 13-15) differ only in dctails of ornamentation, length of preglabellar field, and the slightly more posterior position of palpebral lobes.

Xuanenia Zhou in Zhou et al. 1977, type species X. gramulosa Zhou in Zhou et al. (1977, p. 263, pl. 79, figs. 5-7) from the Linhsiang Formation (Ashgill) of Gaoluo, Xuanen, west Hubei, China, apparently differs little from Prionocleilus. The anterior border is slightly less sharply defined, the front of the glabella is less broadly rounded, and the $1 p$ and $2 p$ lobes appear to coalesce to form composite structures bounded adaxially by longitudinal furrows. The eyes and eye ridges are situated opposite the $2 p$ lobes and furrows, and the librigena shows a row of slim, ventrally directed spines. A possible trace of bacculae is visible in Zhou's pl. 79 , fig. 5 (the holotype) but not in his pl. 79, fig. 6 ; similar structures are seen also in $X$. splendida Ju in Qui et al. (1983, p. 251, pl. 87, figs. 9 and 10), from the Huangnekang Formation (low Ashgill) of Jiande, west Zhejiang, China.

## Family phillipsinellidae Whittington 1950

Genus paraphillipsinella Lu in Lu and Chang 1974
Type species. Paraplillipsinella globosa Lu in Lu and Chang 1974, Pagoda Formation (Caradoc), Chenkou, Sichuan Province, China.

Junior subjective synonym. Protoplillipsinella Chen in Li et al. (1975, p. 155).
Discussion. A translation of the original generic diagnosis, together with minor emendations, was given by Zhou and Dean (1986, p. 766), who followed Lu and Zhou (1981, p. 14) in considering Paraphillipsinella to include, at that time, only two species: P. globosa Lu in Lu and Chang (1974, p. 133, pl. 53, figs. 8 and 9) and $P$. nanjiangensis Lu in Lu and Chang (1974, p. 133, pl. 53,
fig. 10). Of the two, only $P$. globosa is regarded as sufficiently close to the new species to merit discussion, the glabella of $P$. nanjiangeusis being easily recognized by its transversely suboval anterior lobe, short, squat posterior lobe, and wide fixigenae. To the previous generic criteria we now add the presence, sometimes weakly developed, of a narrow anterior border and border furrow immediately adjacent to the axial furrows. A median occipital tubercle is at least sometimes, and possibly always, present.

## Parapliillipsinella pilula sp. nov.

Plate 60, figs. 7, 17-21, 23
Diagnosis. Paraplillipsinella species characterized by: narrow (exsag.) anterior border and border furrow well developed laterally, die out frontally; interocular portion of each fixigena narrow, little more than half width of adjacent part of posterior lobe of glabella; well-defined palpebral lobes sited opposite $2 p$ and posterior half of $3 p$ glabellar lobes.
Holotype. It. 19531 (Pl. 60, fig. 17).
Paratypes. It. 19530 (Pl. 60, fig. 7), It. 19532 (Pl. 60, fig. 18), 1t. 19533 (Pl. 60, fig. 19), It. 19534 (Pl. 60, fig. 20), It. 19535 (Pl. 60, fig. 21), It. 19536 (Pl. 60, fig. 23).
Localities. Şort Dere, Z. 34 and Z.36.
Description and discussion. Rather than give a detailed description of the specics, comments are confined mainly to features relevant to its recognition. Lu's illustrations of the holotype of P. glohosa show the subspherical anterior lobe and subcylindrical posterior lobe of the glabella occupying, respectively, 0.55 and 0.28 of the overall length of the cranidium. In $P$. pilula the corrcsponding figures arc 0.52 and 0.39 in the largest apparently undistorted cranidium but 0.47 and 0.42 in the smallest. In most published illustrations of Paraplillipsinella the boundary between anterior and posterior lobes often appcars as a sharp, transverse furrow; but this is the result of compression and is not secn in spccimens preserved in limestonc (Zhou and Dcan 1986, pl. 62, figs. 13-15). The anterior border and border furrow of the new species arc particularly striking, with each end cxtending adaxially from the shallow axial furrows around the abaxial quarter of the anterior lobe. The posterior lobe is gently tapered and there are four pairs of glabellar lobes, separated by pit-like glabellar furrows; 1 p lobes are slightly larger than $2 p$ to $4 p$ pairs and form part of weakly defined basal glabellar segment. Straight eye ridges run from antcrior ends of palpcbral lobes to meet axial furrows opposite 4 p glabellar lobes. The position of the palpebral lobes in $P$. globosa is difficult to distinguish in published illustrations ( Lu in Lu and Chang 1974, pl. 53, figs. 8 and 9) but is probably opposite the 2 p glabellar lobes, as in the new species. The samc illustration of $P$. globosa showed no anterior extension of the fixigenae beside, and overhung by, the anterior lobe. A whole exoskeleton identificd by Ju (in Qiu et al. 1983) as $P$. lubeiensis Zhou (1974, p. 228, pl. 76, fig. 9) has since been put in synonymy with $P$. nanjiangensis Lu in Lu and Chang 1974 by Zhou and Dean (1986, p. 767). The spccimen lacks the antcrior lobe and shows both the large rostral plate and a forwards extension of the fixigenae, which are as wide as the posterior lobe. A median occipital tubcrcle occurs in at least one paratype of $P$. pilula and may be a general feature of Paraphillipsinella, though not clearly visible in all illustrations.
A subconcentric pattern of anastomosing, finc ridges on the anterior lobe of the glabella extends as subparallel, longitudinal ridges on the posterior lobe. On the holotype there is a suggestion of pits in some of the intervening grooves, an ornamentation gencrally resembling that found in Phillipsinella (Bruton 1976). Evidence for similar ridges in other species of Paraphillipsinella is equivocal or absent, but the holotype of P. globosa Lu in Lu and Chang (1974, pl. 53, figs. 8 and 9) shows rows of fine granules, apparently concentrically arranged. The slightly grcater width (tr.) of the glabella at the 1 p lobes in P. globosa also recalls that in some specics of Phillipsinella (sec Bruton 1976 for various illustrations).
Two incomplete pygidia show a general resemblance to the type species of Plillipsinella, P. parabola (Barrandc 1846) from the Ashgill of Bohemia, redescribed by Whittington (1950, p. 559, pl. 75, figs. 4 and 7). One, a latcx cast (Pl. 60, fig. 19), has the axis slightly abraded but there are traccs of threc axial rings; the pleural regions show three segments, separated by distinct rib furrows and carrying well-dcfined pleural furrows. The other, an internal mould (P1. 60, fig. 7), has four segments and there are traces of ornamentation, comprising oblique, anastomosing ridgcs, similar to that figured by Bruton (1976, pl. 106, fig. 1; pl. 108, figs. 1, 5, 10, 12). Both Turkish pygidia have the posterior margin slightly indented medially and a marginal rim is weakly developed.

## Family asaphidae Burmeister 1843

Subfamily asaphinae Burmeister 1843
Genus birmanites Sheng 1934
Type species. Ogygites birmanicus Reed 1915, from the Hwe Mawng Beds (Lower Ordovician), Hwe Mawng and Hpakhi, northern Shan States, Burma. The synonymy of the genus was discussed by Zhou et al. (1984, p. 17).

## Birmanites latus (Angelin 1851)

Plate 61, figs. 3-7; Plate 62, figs. 1, 2, 4, 8
1851 Niobe lata Angelin, p. 14, pl. 10.
1960 Opsimasaplus latus (Angelin); Kielan, p. 78, pl. 6, figs. 1 and 2; pl. 7, fig. 3; pl. 8, fig. 4; textfig. 20.
1981 Opsimasaplus; Dean, Monod and Perinçek, p. 277.
Figured specimens. It. 19545 (Pl. 61, fig. 3), It. 19546 (Pl. 61, fig. 4), It. 19547 (Pl. 61, fig. 5), It. 19548 (Pl. 61, fig. 6), It. 19549 (Pl. 61, fig. 7), It. 19552 (Pl. 62, fig. 1), It. 19553 (Pl. 62, fig. 2), It. 19554 (Pl. 62, fig. 4), It. 19555 (Pl. 62, fig. 8).

Localities. Şort Tepe, Z.33-4; Şort Dere, Z. 34.
Description and discussion. Birmanites latus is easily the most abundant trilobite in the collections from the Şort Tepe Formation. The matcrial agrees closely with the lectotype and other specimens from the Red Tretaspis Mudstones (Ashgill) of Västergötland, Sweden, described by Kielan (see synonymy above) and provides little additional information.
Some compressed cranidia ( Pl .61 , figs. 4 and 7) appcar to show strong sigmoidal ridges extending from the rear ends of the palpebral lobes, subparallel to the posterior branches of the facial suture, and ending about half-way to the posterior margin. These structures are the result of crushing and are not invariably developed. As in the lectotype cranidium, a large, low, median tubercle is sited just behind a line joining the rear ends of the palpebral lobes. The flanks of the tubercle carry about five or six narrow, subconcentric ridges and the apex has a trace of a small median perforation.

All the pygidia are dorsally compressed with median length slightly more or less than 0.6 of the breadth. The almost straight-sided axis has a frontal breadth about $0 \cdot 2$ that of the pygidium and occupies about 0.8 of its length, though the terminal piece is not well defined. Largest examples (Pl. 61, fig. 3) have at least eight, low, transversely straight axial rings, separated by shallow ring furrows, in the anterior four-fifths of the axis, the remainder being indiscernible. This matches closely Kielan's illustration (1960, pl. 6, fig. 2), where a further two rings and a tiny terminal piece are visible. The pleural fields show, in addition to the large anterior half-ribs, five pairs of ribs clearly defined and a sixth less so; this agrees with the original of Kielan 1960 , pl. 6, fig. 1, though a better-preserved Swedish example (Kielan 1960, pl. 6, fig. 2) has six pairs of ribs and a less well-defined seventh pair.

Kielan did not describe the hypostoma of $B$. latus but two associated Turkish specimens (Pl. 62, figs. 1, 2) of asaphinid type are assigned to the species. Maximum breadth (including anterior wings) is about threequarters the overall length, and the posterior margin is deeply indented to form a narrow, median notch with subparallel sides. Middle body is longitudinally subelliptical with length two-thirds that of hypostoma, and with curved posterior margin concave rearwards, subparallel to median notch; posterolateral extremities

## EXPLANATION OF PLATE 61

Figs. 1 and 2. Diacanthaspis sp. Loc. Z.34. 1, cranidium, 1t. 19542, $\times 6$. 2, left librigena, It. 19543, $\times 6$.
Figs. 3-7. Birmanites latus (Angelin 1851). Loc. Z.34. 3, pygidium, It. 19545, $\times 1 \cdot 5$. 4, cephalon, It. 19546,
$\times 2$. 5, ventral surface of pygidial doublure, It. 19547, $\times 5.6$, exfoliated pygidium showing doublure, It. $19548, \times 2.7$, cranidium showing median tubercle, It. $19549, \times 2 \cdot 5$.
Figs. 8-10. Amphitryon? sp. 8 and 9, loc. Z.34; 10, loc. Z.36. 8, cranidium, It. 19550, $\times 4.9$, cranidium, It. 19551, $\times 5$. 10, front of cranidium showing anterior border, It. 19476, $\times 8$.

,


2


4


6


9



8
8


10
formed by pair of large maculae, bounded anterolaterally by decp, triangular furrows. Posterior borders are large, their lateral margins strongly curved, abaxially convex, linked by low ridges to anterior half of middle body. Frontal portion of hypostoma formed by flat anterior border that circumscribes middle body and widens (exsag.) distally to end in pair of short (tr.), obtusely angular anterior wings. Overall breadth across anterior wings slightly less than that across postcrior borders, and the two structures are separated by broad (exsag.) lateral notches. Except for a few terrace lines around margin of posterior notch and on front of middle body, the surface is smooth.

The hypostoma of B. birmanicus has not been illustrated but that of B. lupeiensis Yi 1957, from the Shihtzupu Formation (Llandeilo) of Guizhou Province, China, was redescribed by Zhou et al. (1984, p. 17, fig. $3 f$ ). It differs from those attributed here to B. latus in having the middle body proportionately shorter and less elliptical in outline; posterior wings are longer and more pointed; median notch is conspicuously wider and longer ( 0.36 versus 0.25 of overall length of hypostoma), its sides converging forwards at $45^{\circ}$ instead of being subparallel; and the anterior border, though not clearly visible, appears to be proportionately shorter.

Family aulacopleuridae Angelin 1854
Subfamily aulacopleurinae Angelin 1854
Genus harpidella M•Coy 1849
Type species. Harpes? megalops M‘Coy 1846, Upper Llandovery of Boocaun, Cong, County Galway, Ireland.

## Harpidella sp.

Plate 60, fig. 22: Plate 62, fig. 6
Figured specimens. It. 19541 (Pl. 60, fig. 22), It. 19559 (Pl. 62, fig. 6).
Locality. Şort Dcre, Z.34.
Description and discussion. This form is assigned to Harpidella on account of the large, posteriorly situated palpebral lobes, considered by Thomas and Owens (1978, p. 71) as an important character in distinguishing Harpidella from Otarion Zenker 1833. However, it also resembles the type species of Otarion, O. diffractum Zenker 1833 (see Thomas and Owens 1978, pl. 7, figs. 1-3, 5, 6), from the Kopanina Formation (Ludlow) of Dlouhá Hora, Czechoslovakia, in the narrow glabella, narrow (sag.) anterior border, faint palpebral furrows, and vaulted preglabellar field. The Turkish specimens are inadequate for satisfactory comparison but are generally similar to undetermined species of Otarion figured by Ingham (1970, pl. 5, fig. 12) from the Ashgill, Cautleyan Stage, of northern England and by Dean (1974, pl. 26, fig. 9) from the Chair of Kildare Limestone (Ashgill, Rawtheyan) in eastern Ireland, especially in the outline and size of the glabella, and in the small Ip glabellar lobes. Both these British and Irish species have large, backwardly placed palpebral lobes and are probably better referred to Harpidella.

## EXPLANATION OF PLATE 62

Figs. 1, 2, 4, 8. Birmanites latus (Angelin 1851). 1, 2, 4, loc. Z.34; 8, loc. Z.33-4. 1, hypostoma, It. 19552, $\times 3$. 2, hypostoma, It. 19553, $\times 3.4$, right librigena, It. 19554, $\times 2.8$, dorsal exoskeleton, It 19555, $\times 3$.
Fig. 3. Miraspis sp. Loc. Z.34. Cranidium, It. 19556, $\times 7.5$.
Fig. 5. Stenopareia sp. Loc. Z.36. Cranidium, It. 19558, $\times 5$.
Fig. 6. Harpidella sp. Loc. Z.34. Cranidium, It. 19559, $\times 9$.
Fig. 7. Genus and spccies undetermined. Loc. Z.34. Pygidium, It. 16062, $\times 3.5$.
Figs. 9? and 13. Dicranopeltis sp. Loc. Z.34. 9, right side of cranidium, It. 19471, $\times 3 \cdot 5$. 13, pygidium, It. 19472, $\times 4$.
Fig. 10. Diacallthaspis sp. Loc. Z.34. Pygidium, It. 19557, $\times 7$.
Figs. 11?, 12, 14. Lichas aff. laciniatus (Wahlenberg 1821). Loc. Z.34. 11, fragment of cranidium, 1t. 19473, $\times 4.12$, cranidium, It. $19474, \times 3.14$, cranidium, It. $19475, \times 3$.


Type species. Phoroceplala typa Lu in Lu et al. 1965, Siliangssu Formation (upper Arenig), Laingshan, south Shaanxi Province, China.

## Phorocephala sp.

Plate 60, figs. 4 and 24

Figured specimens. It. 19528 (Pl. 60, fig. 4), It. 19529 (Pl. 60, fig. 24).
Locality. Şort Dere, Z.34.
Description. Cranidium about twice as long as wide. Glabella well defined by deeply ineised axial furrows, its length 0.65 that of eranidium, slightly longer than wide, gently tapered forwards, rounded anteriorly. Oecipital ring one-sixth the cranidial length (sag.), wider (tr.) than base of glabella, with posterior margin arched backwards; abaxial portions are narrower (exsag.) and curve forwards slightly to axial furrows. Fixigenae narrow, with width one-sixth that of cranidium as measured aeross mid-length of palpebral lobes. Palpebral lobes gently eurved in plan and run slightly inwards anteriorly; their length is 0.45 that of eranidium and they extend almost to posterior border furrow. Anterior branehes of faeial suture subparallel; preglabellar area short, equal to one-sixth the cranidial length. Anterior border is dorsally eonvex, widens adaxially, and is defined by distinet, though shallow, anterior border furrow. Preglabellar field depressed, as long (sag.) as anterior border.
The pygidium is about twiee as broad as long, its outline approximately lozenge shaped. The large, strongly tapered axis has a frontal breadth half that of the pygidium. There are two large, curved axial rings, convex forwards, with traees of a third; the pleural regions have a very thin, marginal rim and show two pairs of deep pleural furrows and two pairs of shallow rib furrows.
Discussion. The cranidium, though incomplete, is comparable with that of the type species, P. typa Lu (in Lu et al. 1965, p. 587, pl. 123, fig. 14; see also Lu 1975, pl. 34, fig. 13), in the shape of the glabella and the size and location of the palpebral lobes. According to Zhou and Dean (1986, p. 751) the preglabellar field of Phorocephala is absent in adult cranidia, though present in immature cranidia, of all known species of Caradoc and Ashgill age. The present specimen has a median length of only 2.3 mm and probably represents a juvenile individual.

The pygidium of the type species was not described by Lu but that of P. shizipuensis Yin (in Yin and Lee 1978) from the Llandeilo of Guizhou Province, China, figured by Zhou et al. (1984, fig. $5 x, z$ ) has an outline resembling that of the present specimen. Both have a large, triangular axis but in the Turkish species the pleural regions are proportionately smaller, with straighter margins, and there are only two pairs of pleural and interpleural furrows, compared with four.

Family remopleurididae Hawle and Corda 1847
Genus amphitryon Hawle and Corda 1847
Type species. Caphyra radians Barrande 1846, p. 32 (a senior subjective synonym of Caphyra murchisonii Hawle and Corda 1847), from the Králův Dvůr Formation (Ashgill), Králův Dvỉr, Czeehoslovakia.

## Amphitryon? sp.

Plate 61, figs. 8-10
Figured specimens. It. 19550 (Pl. 61, fig. 8), 1t. 19551 (Pl. 61, fig. 9), It. 19476 (Pl. 61, fig. 10).
Localities. Şort Dere, Z. 34 and Z. 36 .
Description and discussion. The most complete cranidium, though dorsally compressed, is apparently of low convexity with overall breadth of 9.0 mm and median length (including preglabellar field) of 9.8 mm . Cranidial length, excluding anterior tongue of glabella, is 7.2 mm , and basal breadth of glabellar tongue is 0.3 of maximum glabellar breadth. Glabellar outline closely resembles that of Amphitryon radians and the most
obvious difference is the total absence of glabellar furrows in the Turkish material, though the value of this as a generic character is unknown. Whittington's (1966, p. 72, text-fig. $4 a-g$ ) illustrations of $A$. radians from Bohemia show three incised pairs of glabellar furrows, but the occipital ring and palpebral lobes are virtually indistinguishable from those of the Turkish specimen, and the glabellar tongue ends behind a triangular preglabellar field, the apex of which probably coincided with a median suture. The breadth (tr.) of the glabellar tongue as shown by Whittington equals only about 0.14 the glabellar breadth, but in material from Bohemia and Poland assigned to $A$. radians by Kiclan (1960, pl. 2, figs. 3, 5, 6) the corresponding figure varies from 0.2 to 0.25 , though a small Polish cranidium identified as Amphitryon sp. (Kielan 1960, pl. 3, fig. 12) has a very narrow glabellar tongue (only 0.13 estimated). It is possible that the relative breadth of the glabellar tongue changed during ontogeny, and for present purposes more importance is attached to the triangular preglabellar field, which readily distinguishcs the Turkish specimens from Remopleurides (see account of type species $R$. colbii Portlock 1843 in Whittington 1950, p. 540).

Other species previously assigned to Remopleurides that have a triangular preglabellar ficld include the Chinese forms R. nasutus Lu 1957 (see Lu 1975, pp. 109, 299, pl. 3, fig. 15; pl. 4, figs. 5, 8, 9), from the Arenig of south Shensi and West Hupeh, and R. shiltzupuensis Lu 1957 (see Lu 1975, pp. 111, 301, pl. 4, fig. 15), from the Arenig to Llandeilo of North Kueichou and West Hupeh. R. nasutus has a wide glabellar tongue with rounded frontal margin and the triangular preglabellar field is shorter and less distinct than in the Turkish material; R. shilhzupuensis is closer to the latter in the shape of the preglabellar field but has a much longer glabellar tongue. Sculptella and Sculptaspis from the Middle Ordovician of Norway (Nikolaisen 1982, pp. 265, 276) bear a superficial resemblance to the present material but their cranidia lack the triangular preglabellar field.

Family illaenidae Hawle and Corda 1847 Genus stenopareia Holm 1886

Type species. Illaenus Linnarssoni Holm 1882, Boda Limestonc (Ashgill Series), Dalarne, Sweden.
Stenopareia sp.
Plate 62, fig. 5
Figured specinen. It. 19558.
Locality. Şort Dere, Z. 36 .
Description and discussion. The Turkish cranidium is too compressed for specific identification, but the form of the axial furrows and the width of the fixigenae are consistent with those of Stenopareia limnarssoni, redescribed by Warburg (1925, p. 117, pl. 2, figs. 14 18), though the position of the eyes is not clear.

Family lichidae Hawle and Corda 1847
Subfamily lichinae Hawle and Corda 1847
Genus lichas Dalman 1827
Type species. Entomostracites laciniatus Wahlenberg 1821, from the Dalmanitina Beds (Ashgill) of Bestorp, Mösseberg, Sweden.

## Lichas aff. laciniatus (Wahlenberg 1821)

Plate 62, figs. 11?, 12, 14
Figured specimens. It. 19473 (Pl. 62, fig. 11), It. 19474 (Pl. 62, fig. 12), It. 19475 (Pl. 62, fig. 14).
Locality. Şort Dere Z.34.
Description and discussion. Two cranidia, slightly distorted by dorsal compression, show clearly the composite lateral lobes incompletely defined posterolaterally as in Lichas laciniatus, redescribed by Warburg (1925, p. 295 , pl. 8 , figs. $16,17,20 ; 1939$, p. 15 , pl. 9 , fig. $3 a, b$ ). The occipital lobes are also similar and the palpebral lobes correspond in size and location. The long axes of the composite lateral lobes diverge forwards at about 45 , comparable with Warburg's illustrations, but the median lobe narrows to half the breadth of the composite lobes; this contrasts with $L$. laciniatus, where the median and composite lobes are of equal breadth
in larger specimens, though one small cranidium (Warburg 1925, pl. 8, fig. 17) has the median lobe slightly narrower. The narrowest part of the median lobe is also much narrower than that of the neotype of L. affinis Angelin, 1854, from the Ashgill of Sweden (Warburg 1939, pl. 9, fig. 13). The glabella in the Turkish material is relatively shorter than that of L. laciniatus (length : breadth $=34: 34$ versus $38: 34$ ), the occipital ring is proportionately narrower (exsag.), especially distally, and the occipital lobes are notably larger. L. laciniatus ranges from the Ashgill into the Llandovery Series and cranidia from northern England described by Temple (1969) correspond to the Swedish material.

## Genus dicranopeltis Hawle and Corda 1847

Type species. Lichas scabra Beyrich 1845, upper part of the Liteň Formation (Wenlock), Svaty Jan, near Beroun, Czechoslovakia.

## Dicranopeltis sp.

Plate 62, figs. 9? and 13
Figured specimens. It. 19471 (Pl. 62, fig. 9), It. 19472 (Pl. 62, fig. 13).
Locality. Şort Dere, Z. 34 .
Description and discussion. The Turkish pygidium generally resembles material from the Ashgill (Boda Limestonc and Dalmanitina Beds) of Sweden described by Warburg, first as Dicranopeltis elegans (Törnquist 1884) (Warburg 1925 , p. 291, pl. 7, figs. 27 and 31 ; pl. 8, figs. 9 and 10), and later put in synonymy with $D$. polytomus (Angelin 1854) (Warburg 1939, p. 134, pl. 11, figs. 4-6). D. sp. differs from the pygidium of $D$. polytonus in having a less sharply constricted terminal piece, and the axial furrows are moderately convergent to what is essentially a low, slightly tapered post-axial ridge, separated by a pair of very shallow grooves from the extremities of the third pleurae, which end in short free points separated by a very small median notch. The third pleurae of the Turkish pygidium occupy a relatively smaller area and the equisized pleural bands end in rounded tips at the broad border. The entire surface, except furrows, is covered with closely spaced, small tubercles that become slightly smaller on the posterior border. The rearmost part of the axis, immediately in front of the post-axial ridge, is bulbous as in the Swedish material.

The sole, possibly associated cranidial fragment is crushed and the position of some furrows apparently displaced, so that the composite lateral lobes appear more divergent forwards than originally, and the palpebral lobe has been displaced towards the glabella. The occipital ring and right occipital lobe resemble those of $D$. polytonms, but the rearmost part of the axial furrow is almost indiscernible, in marked contrast to the well-defined furrow in the Swedish species.

## Family odontopleuridae Burmeister 1843

Subfamily odontopleurinae Burmeister 1843
Genus diacanthaspis Whittington 1941
Type species. Diacantlaspis cooperi Whittington 1941, Lower Martinsburg Formation (Caradoc), Virginia, USA.

## Diacantlaspis sp.

Platc 61, figs. 1 and 2; Plate 62, fig. 10
Figured specimens. It. 19542 (Pl. 61, fig. 1), It. 19543 (Pl. 61, fig. 2), It. 19557 (Pl. 62, fig. 10).
Locality. Şort Dere, Z. 34 .
Description and discussion. Cranidium has parallel-sided glabella with wide, subparallel-sided median glabcllar lobe and narrow, clongated 1 p and 2 p lateral glabellar lobes. It agrecs well with the cranidium of $D$. sp. of Lu and Zhou (1981, p. 20, pl. 3, fig. 10) from the Tangtou Formation (low Ashgill) of the Nanjing Hills, China. In each case the cranidium is incomplete, and no associatcd pygidium is available for the Chinese form, so that specific identity cannot be established. Occipital spines are not preserved on the Turkish spccimen, though a prominent median tubercle is present close to the occipital furrow. Another spccics of Diacanthaspis with parallcl-sided glabella is D. laokuangshanensis Lu and Chang (1974, p. 136, pl. 56, figs.

5-7), from the Wufeng Formation (Ashgill) of western Sichuan, China, but the lateral glabellar lobes are wider (tr.) and the median lobe expands forwards.

A fragmentary Turkish librigena shows a flat, coarsely granulate genal field and small, raised eye socle. Lateral border is narrow, ridge-like, with marginal spines that become successively shorter (tr.) anteriorly.

The pygidium has a semicircular outline and a convex axis comprises two axial rings and short terminal piece. Pleural field narrow (tr.), weakly defined by narrow border of low convexity. There are five equispaccd pairs of slender, radiating border spincs, whereas in all other known species of Diacanthaspis there are six or more pairs.

## Subfamily miraspidinae R. and E. Richter 1917 <br> Genus miraspis R. and E. Richter 1917

Type species. Odontoplenra mira Barrande 1846, from the Liteň Formation (Wenlock), Lodenice, Czechoslovakia.

## Miraspis sp.

Plate 62, fig. 3
Figured specimen. It. 19556.
Locality. Şort Dere, Z. 34 .
Description. Glabella tapers gently forwards, its basal breadth half that of cranidium. It is transversely convex with elongated, subrectangular median lobe that expands abruptly to form very short frontal glabellar lobe. Three pairs of lateral glabellar lobes well defined by deep longitudinal furrows; Ip lobes oval in outline, their width equal almost to that of median glabcllar lobe and to 0.4 of glabellar length; 2 p lobes sub-square in plan, slightly narrower than Ip lobes; 3p lobes tiny, transverse, with length (exsag.) one-quarter that of 2 p lobes. Three pairs of almost transverse lateral glabellar furrows present; 1 p and 2 p furrows deeply incised, 3p pair shallow. Occipital furrow shallow, broad; anterior part of occipital ring with median tubercle and pair of long, broadly based spines that extend upwards posterolaterally; posterior part of occipital ring not preserved. Axial furrows distinct posteriorly, shallow anteriorly. Fixigenae narrower posteriorly than lp lobes, and become still narrower further forwards. Palpebral lobes situated opposite mid-points of 1 p lobes. Sutural and palpebral ridges slightly convex abaxially and converge forwards. Anterior border as narrow as palpebral ridges, and defined by shallow preglabellar furrow. Cranidial surface covered with densely spaced tubercles of different sizes.

Discussion. The cranidium, though imperfectly preserved, is compatible with that of the lectotype of Miraspis mira (Barrande 1852, pl. 39, fig. 3, selected by Přibyl in Horný and Bastl 1970, p. 203) and of a well-preserved dorsal exoskeleton of that species figured by Prantl and Vaněk (in Horný et al. 1958, pl. 5, fig. 1) but differs in having the median glabellar lobe narrower, with straighter sides, while the 1 p and 2 p glabellar furrows are more distinct abaxially.

Ordovician species of Miraspis have been recorded from Sweden (Whittington and Bohlin 1958; Bruton 1966), Norway (Owen and Bruton 1980), eastern Ireland (Dean 1974), North Wales and Scotland (Whittington and Bohlin 1958, p. 43). In the shape of the cranidium and median glabellar lobe and in the pattern of surface granulation the Turkish specimen resembles $M$. sp. of Owen and Bruton (1980, p. 36, pl. 10, figs. 18 and 20), from the uppermost Solvang Formation (probably low Ashgill) of Ringerike, Norway, but the latter has shallower 1 p and 2 p glabellar furrows and the 2 p glabellar lobes are proportionately much smaller.

# Genus and species undetermined 

Plate 62, fig. 7
Figured specimen. It. 16062.
Locality. Şort Dcre, Z. 34 .
Description and discussion. A fragmentary internal mould, interpreted tentatively as part of a pygidium, has the surface, excluding furrows, covered with coarse, closely spaced tubercles. The ornamentation bears some
resemblance to that of the odontopleurids, and there is some evidence of one of a pair of subparallel ridges ending in line with the first axial ring; in the odontopleurids similar ridges link the corresponding ring with a pair of large marginal spines. Three markedly unequal axial rings, separated by transversely straight ring furrows that deepen abaxially, arc followed by a minute, weakly defined terminal piece and gently declined postaxial field. No satisfactory comparison was made.

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## REFERENCES

altinli, i. e. 1963. Explanatory text of the geological map of Turkey, Cizre. Maden Tetkik Arama Enstit. Haritasi, 103 pp .
angelin, n. p. 1851-1878. Palaeontologia Scandinavica: Academiae Regiae Scientiarum Suecanae (Holmiae); Pars I. Crustacea formationis transitionis, pp. 1-24, pls. 1-24 [1851]; Pars II, pp. i-ix, 25-92, pls. 25-41 [1854]; republished in revised and combined form (ed. G. Lindström), pp. x + 96, pls. 1-42 [1878].
barrande, J. 1846. Notice préliminaire sur le Système silurien et les Trilobites de Bohême, vi +97 pp . Leipzig.
_— 1852. Système silurien du centre de la Bohême. lère partie. Recherches paléontologiques. Vol. I. Crustacés,
Trilobites, $\mathrm{xxx}+935 \mathrm{pp}$. Prague and Paris.
beyrich, e. 1845. Ueber einige bölmischen Trilobiten, 47 pp. Reimer, Berlin.
bruton, D. L. 1966. A revision of the Swedish Ordovician Odontoplcuridae (Trilobita). Bull. geol. Instm Univ. Uppsala, 43, 1-40.

- 1968. A revision of the Odontopleuridae (Trilobita) from the Palaeozoic of Bohemia. Skr. norske Vidensk-Akad. 25, 73 pp.
- 1976. The trilobite genus Phillipsinella from the Ordovician of Scandinavia and Great Britain. Palaeontology, 19, 699-718.
burmeister, h. 1843. Die Organisation der Trilobiten, 147 pp. Berlin.
campbell, k. s. w. 1967. Trilobites of the Henryhouse Formation (Silurian) in Oklahoma. Butl. Okla. geol. Surv. 123, 1-227.
chugaeva, m. n. 1958. The Ordovician of Kazakhstan. III: The Ordovician trilobites of the Chu-Ili Mountains. Trudy geol. Inst., Leningr. 9, 5-138. [In Russian.]
dalman, J. w. 1827. Om Palaeaderna, eller de så kallade Trilobiterna. Bih. K. Svenska VetenskAkad. Handl. 1, 113-152, 226-294.
dean, w. t. 1959. Duftonia, a new trilobite genus from the Ordovician of England and Wales. Palaeontology, 2, 143-149.
- 1966. The Lower Ordovician stratigraphy and trilobites of the Landeyran Valley and the neighbouring district of the Montagne Noire, south-western France. Bull. Br. Mus. nat. Hist. (Geol.), 12, 245-353.
- 1967a. The distribution of Ordovician shelly faunas in the Tethyan region. In Adams, C. J. and ager, D. v. (eds.). Aspects of Tethyan biogeography. Systematics Assoc. Spec. Publ. 7, 11-44.

1967b. The correlation and trilobite fauna of the Bedinan Formation (Ordovician) in south-eastern Turkey. Bull. Br. Mus. nat. Hist. (Geol.), 15, 81-123.

- 1971. The trilobites of the Chair of Kildarc Limestone (Upper Ordovician) of eastern Ireland. Palaeontogr. Soc. [Monogr.], part 1, 1-60.
- 1974. 1bid. 2, 61-98.
- 1980. The Ordovician System in the Near and Middle East. Correlation chart and explanatory notes. Publ. int. Union geol. Sci. 2, 1-22.
- MONOD, O. and PERINÇEK, D. 1981. Correlation of Cambrian and Ordovician rocks in southeastern Turkey. Petroleum Activities at the 100th Year (100 Yilda Petrol Faaliycti). Türkiye Cümhuriyet Petrol Isleri Genel Müdürlüğü Dergisi, 25, 269-291 (English), 292-399 (Turkish).
glimberg, c. F. 1961. Middle and Upper Ordovician strata at Lindegard in the Fågclsång district, Scania, S. Swcden. Geol. För. Stockh. Förh. 83, 79-85.
hawle, J. and Corda, A. J. c. 1847. Prodrom einer Monographie der böhmischen Trilobiten, 176 pp. J. G. Calvé, Prague.
holliday, s. 1942. Ordovician trilobites from Nevada. J. Paleont. 16, 471-478.
holm, g. 1882. De Svenska artena af Trilobitslagct Illaenus (Dalman). Bil. K. Svenska Vetensk Akad. Handl. 7 , xiv +148 pp .

1886. Illaeniden. Revision der Ostbaltisehen silurischen Trilobiten von. Fr. Schmidt. Abt. III. Mem. Acad. imp. Sci. St.-Petersb. (7) 33, 1-173.
horný, r. and bastl, f. 1970. Type specimens of fossils in the National Museum, Prague. Volume 1. Trilobita. 354 pp . Museum of Natural History, Prague.
prantle, f. and vaněk, J. 1958. K. otázce hraníce mezi wenlockem a ludlowem v Barrandienu; $S b$. Ústréed. Úst. geol. 24, 217-256 (in Czech), 257-278 (in English).
hupé, P. 1955. Classifieation des Trilobites. Annls Paléont. 41, 91-325 (111-345).
ingham, J. к. 1970. The Upper Ordovician trilobites from the Cautley and Dent distriets of Westmorland and Yorkshire. Palaeontogr. Soc. [Monogr.], part 1, 1-58.

- 1974. Ibid. 2, 59-87.
jaanusson, v. 1963. Classification of the Harjuan (Upper Ordovieian) roeks of the mainland of Sweden. Geol. För. Stockh. Förh. 85, 110-144.
- 1982. Introduction to the Ordovieian of Sweden. In bruton, D. L. and williams, s. h. (eds.). Field excursion Guide. IV. International Symposium on the Ordovician System. Paleont. Contr. Univ. Oslo, 279, 1-217.
Janvier, P., Lethiers, f., monod, o. and balkaş, o. 1984. Discovery of a vertebrate fauna at the Devonian Carboniferous boundary in SE Turkey (Hakkâri Province). Jl Petrol. Geol. 7, 147-168.
jones, t. r. and woodward, h. 1898. A monograph of the British Palaeozoic Phyllopoda (Phyllocarida, Paekard). Palaeontogr. Soc. [Monogr.], part 3, 125-176.
kielan, z. 1960. Upper Ordovician trilobites from Poland and some related forms from Bohemia and Scandinavia. Palaeont. pol. 11, 1-198.
kobayashi, t. 1935. The Cambro-Ordovician formations and faunas of South Chosen. Palaeontology, Pt. 3, Cambrian faunas of South Chosen with special study on the Cambrian trilobite genera and families. J. Fac. Sci. Tokyo Univ. 4, 49-344.
- 1951. On the Ordovieian trilobites in Central China. Ibid. Sect. II, 8, 1-87.
- 1960. Some Ordovician fossils from east Tonkin, Vict Nam. Jap. J. Geol. Geogr. 31, 39-48.
koroleva, m. N. 1959. New species of trilobites from the Middle and Upper Ordovieian of northern Kazakhstan. Dokl. Akad. Nauk SSSR, 124, 1313-I316. [In Russian.]
li yaoxi, song lisheng, zhou zhiolang and yang ingyao. 1975. Stratigraplical Gazetteer of Lower Palaeozoic, western Dabashan, 372 pp. Geological Publishing House, Beijing. [In Chinese.]
linnarsson, J. g. o. 1869. Om Vestergötlands Cambriska och Siluriska Aflagringar. Bil. K. Svenska Vetensk Akad. Handl. 8 (2), 1-89.
lu yanhao. 1975. Ordovician trilobite faunas of eentral and southwestern China. Palaeont. sin. 11, I-484.
——and chang wentang. 1974. Ordovician trilobites, 124-136. In A handbook of Stratigraphy and Palaeontology in southwest China. Seienee Press, Beijing. [In Chinese.]
-- chu chaoling, chien yiyuan and hisiang leewen. 1965. Trilobites of China, 766 pp . Seicnee Press, Beijing. [In Chinese.]
- and zhou zhiyı. 1979. Systematic position and phylogeny of Hammatocnemis (Trilobita). Acta palaeont. $\sin$. 18, 415-434. [In Chincse with English summary.]

198I. Early Upper Ordovieian trilobites from the Nanjing Hills. Bull. Nanjing Inst. Geol. Palaeont. 3, 1-27. [In Chinese with English abstract.]
m'coy, F. 1846. A synopsis of the Sihrian fossils of Ireland, 72 pp . MeGlashand and Gill, Dublin.

- 1849. On the classification of some British fossil Crustacea, with notiees of some new forms in the University Colleetion at Cambridgc. Ann. Mag. nat. Hist. (2) 4, 161-179, 330-335, 392-414.
moore, r. C. (ed.). 1959. Treatise on Invertebrate Paleontology, Pt. O, Arthropoda I, xix +560 pp . Geologieal Soeiety of Amcrica and University of Kansas Press, Lawrence, Kansas.
nicholson, h. a. and etheridge, r. 1878. A monograph of the Silurian fossils of the Girvan District in Ayrshire. I, ix +135 . Edinburgh and London.
nikolaisen, f. 1982. The Middle Ordovician of the Oslo Region, Norway, 32. Trilobites of the Family Remopleurididae. Norsk geol. Tidsskr. 62, 231-329.
olin, e. 1906. Om de Chasmopskalken oeh Trinucleusskiffern motsvarande bildningarne i Skåne. Acta Univ. lund. NF [2], 2 (3), 1-79.
ÖPIK, A. A. 1937. Trilobiten aus Estland. Acta Comment. Univ. tartu. (A) 32 (3), 1.163.

Öplк, A. A. 1967. The Mindyallan Fauna of north-western Quecnsland. Brll. Bhr. Miner. Resour. Geol. Geoplys. Aust. 74, 1-404.
OWEN, A. and bruton, D. L. 1980. Late Caradoc-carly Ashgill trilobites of the central Oslo region, Norway. Paleont. Contr. Univ. Oslo, 245, 63 pp.
PORTLOCK, J. E. 1843. Report on the geology of the county of Londonderry, and of parts of Tyrone and Fcrmanagh, xxi +784 pp. Dublin (Milliken, Hodges and Smith) and London (Longman, Brown, Green, and Longmans).
qiU hongan, lu yanhao, zhu zhaoling, bi dechang. lin tianrui, zhou zhiyi, zhang quanzhong, qian yiyuan, ju tianyin, han nairen and wei xiuzhe 1983. Trilobita. In Palaeontological Atlas of east China. II, 28 -254. Geological Publishing House, Beijing. [In Chinese.]
reed, f. R. C. 1915. Supplementary memoir on new Ordovician and Silurian fossils from the Northern Shan States. Mem. geol. Surv. India Palaeont. indica, ns 6 (1), 1-123.
reynolds, s. h. 1894. Woodwardian Museum Notes. Certain fossils from the Lower Palacozoic rocks of Yorkshire. Geol. Mag. 31, 108-111.
richter, r. and richter, e. 1917. Über dic Einteilung der Familie Acidaspidae und über einige ihrer devonischen Vertreter. Zentbl. Miner. Geol. Paläont. (1917), 462-472.
rouault, m. 1847. Extrait du Mćmoire sur les Trilobites du Département d’Ille-et-Vilaine. Bull. Soc. géol. Fr. (2) 4, 309-328.
sars, m. 1835. Ueber einige neue oder unvollstandig bekannte Trilobiten. Isis, Jcha (1835), 4.
shaw, f. C. and ormiston, a. r. 1964. The eye socle of trilobites. J. Paleont. 38, 1001-1002.
sheng, S. F. 1934. Lower Ordovician trilobitc fauna of Chekiang. Palacont. sin. B 3 (1), 1-19.

- 1974. Ordovician trilobites from western Yunnan and its stratigraphical significance. In Subdivision and correlation of the Ordovician System in China, T-153, 96-140. Geological Publishing House, Beijing. [In Chinese.]
sun, y. c. 1931. Ordovician trilobites of Central and Southern China. Polocont. sin. B 7, 1-47.
temple, J. т. 1954. The hypostome of Encrimurus variolaris and its relation to the cephalon. Geol. Mag. 91, 315-318.
- 1969. Lower Llandovery (Silurian) trilobites from Keisley, Westmorland. Bull. Br. Mus. nat. Hist. (Gcol.), 18, 197-230.
thomas, A. т. 1978. British Wenlock trilobites. Palaeontogr. Soc. [Morogr.], 1-56.
- and owens, r. m. 1978. A review of the trilobite family Aulacopleuridae. Palaeontology, 21, 65-81.
törnquist, s. L. 1884. Undersökningar ofver Siljansområdets trilobitfauna. Sver. geol. Unders. Afh. C 66, 1-101.
TRIPP, R. P. 1958. Stratigraphical and geographical distribution of the named species of the trilobite superfamily Lichacea. J. Paleont. 32, 574-582.
vogdes, A. w. 1890. A bibliography of Palaeozoic Crustacea from 1698 to 1889, including a list of North American species and a systematic arrangement of genera. Bull. US geol. Surv. 63, 1-177.
wahlenberg, G. 1821. Petrificata Telluris Svccanae. Nova Acta R. Soc. Scicht. upsala, 8, 1-116, 293-297.
warburg, e. 1925. The trilobites of the Leptaena Limestone in Dalarne. Bull. geol. Instn Univ. Uppsala, 17, 1446.

1939. The Swedish Ordovician and Lower Silurian Lichidac. Bih. K. svensk Vetensk Akad. Handl. 17 (4), 1-162.
weber, v. n. 1932. Trilobites of Turkestan. Izd. Vses. Geol.-Razv. Ob'ed. NKTP, iv +157 pp. [In Russian with English summary.]
1940. Trilobites of the Silurian beds. No. 1. Lower Silurian trilobites. Monogr. Palaeont. USSR, 69 (1), 1-110. [In Russian.]
whittard, w. f. 1955. The Ordovician trilobites of the Shelve inlier, west Shropshire. Palacontogr. Soc. [Monogr.], part 1, 1-40.
whittington, h. b. 1941. Silicified Trenton trilobites. J. Paleont. 15, 492-522.

- 1950. Sixteen Ordovician genotype trilobites. Ibid. 24, 531-565.
- 1965. A monograph of the Ordovician trilobites from the Bala area, Merioneth. Palaeontogr. Soc. [Monogr.], part 2, 33-62.
- 1966. Ibid. 3, 63-92.
- and bohlin, b. 1958. New Lower Ordovician Odontopleuridae (Trilobita) from Öland. Bull. geol. Instn Univ. Uppsala, 38, 37-45.
williams, a., Strachan, i., bassett, d. a., dean, w. t., ingham, J. k., wright, a. d. and whittington, H. B. 1972. A correlation of Ordovician rocks in the British Isles. Geol. Soc. Lond., Spcc. Rep. 3, 1-74.
yi yongen. 1957. The Caradocian trilobite fauna from the Yangtze Gorgcs. Acta Palaeont. sin. 5 (4), 527 560. [In Chinese with English summary.]
yin gongzheng and lee shansi. 1978. Trilobita. In Palaeontological Atlas of Southwest China. Guizhou Province (1), 385-595. Geological Publishing House, Beijing. [In Chincse.]
zenker, J. C. 1833. Beitrage zur Naturgeschichte der Utwelt, Organische Reste (Petrefacten) aus der Altenhurger Braunkohlen-Formation, dem Blankenburger Quadersandstein, jenaischen bumten Sandstein und böhmischen Uebergangsgebirge, viii +67 pp . Jena.
zhang tairong. 1981. Trilobita. In Atlas of palacontology of NW China: Xinjiang Volume 1, 134-213. Geological Publishing House, Bcijing. [In Chinese.]
zhang wentang. 1979. On the Miomera and Polymera (Trilobita). Scientia sim. 10, 996 -1004. [In Chinese.] - li jijin, ge meiyu and chen junyuan. 1982. Subdivision and corrclation of the Ordovician in ChinaCorrelation chart and its explanation for the Chinese Ordovician System. In Correlation charts and their explanations for Chinese strata, 55 72. Science Press, Bcijing.
zhou tianmei, liu yiren, meng xiansong and sun zhenhua. 1977. Trilobita. In Atlas of Palacontology of central and south China, I40-266. Gcological Publishing House, Beijing. [In Chinese.]
zhou zhiyi and dean, w. t. 1986. Ordovician trilobites from Chedao, Gansu Province, north-west China. Palaeontology, 29, 743-786.
yin gongzheng and tripp, r. p. 1984. Trilobitcs from the Ordovician Shihtzupu Formation, Zunyi, Guizhou Province, China. Trans. R. Soc. Edinb. 75, 1336.

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